

# Non-opaque and Opaque Ice Cloud Properties from Infrared Radiances at 3.7, 6.7, 11.0, and 12.0 $\mu\text{m}$

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## IR Retrieval?

- weather and climate modeling
  - cloud diurnal cycle
  - cloud climatology, long term variation
- is still remained the challenge
  - Not like VIS, IR radiances of clouds are more sensitive to cloud and atmospheric profiles
  - IR blackbody limitation, effective emissivity close to 1, reliable  $\tau < 6-8$

## Only 3.7, 6.7, 11.0, and 12.0- $\mu\text{m}$ bands?

- focus on retrieval algorithms using fewer IR bands, mostly common.
- cloud retrieval algorithms utilized as few channels that are common to most meteorological satellite imagers as possible have the advantage of producing consistent cloud properties (Minnis et al., 2011)

## Objectives

- physical retrievals of cloud-top temperature/height, effective emissivity,  $\tau$  for non-opaque ice clouds
- neural network training for opaque ice cloud  $\tau$

# Estimating Non-opaque Ice Cloud Tc from 6.7 and 11.0-μm Radiances

Current CERES nighttime cloud retrieval algorithm:  
The Shortwave-infrared Infrared Split-window Technique (SIST) for all surfaces at night (CERES, Minnis et al., 1998; 2011) using 3.7, 11.0, and 12.0 μm.

When there is no 3.7 or 12 μm available (e.g., some GEOs)

Szejwach (1982)

$$I_{WV} = (1 - N\epsilon_{WV}) \cdot I_{WV}^{clear} + N\epsilon_{WV} \cdot I_{WV}^{cloud}$$

$$I_{11} = (1 - N\epsilon_{11}) \cdot I_{11}^{clear} + N\epsilon_{11} \cdot I_{11}^{cloud}$$

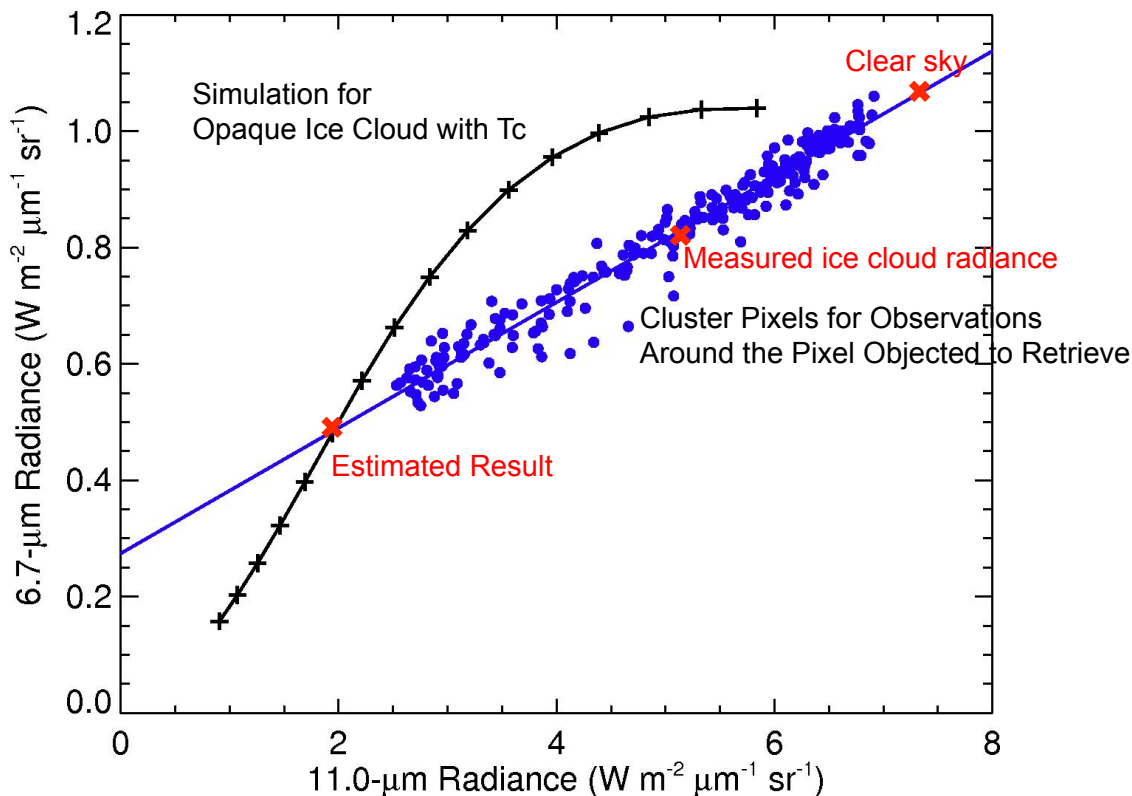
$$\frac{I_{WV} - I_{WV}^{clear}}{I_{11} - I_{11}^{clear}} = \frac{N\epsilon_{WV} (I_{WV}^{cloud} - I_{WV}^{clear})}{N\epsilon_{11} (I_{11}^{cloud} - I_{11}^{clear})}$$

Clear Sky Radiance

Satellite Measured Radiance

Effective Emissivity

Cloud Radiance at Tc Blackbody



## Pixel-based:

Tc is derived by matching the ratio along black line

Could be strongly affected by variation of clear sky radiance

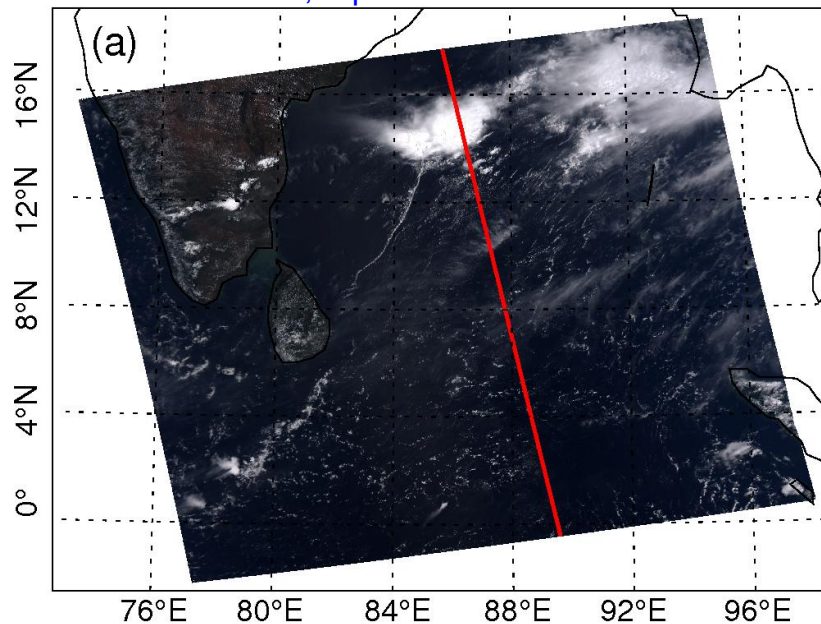
## Cluster-based:

Linear fitting cluster pixels, crossing with simulated opaque ice cloud ratio (black line)

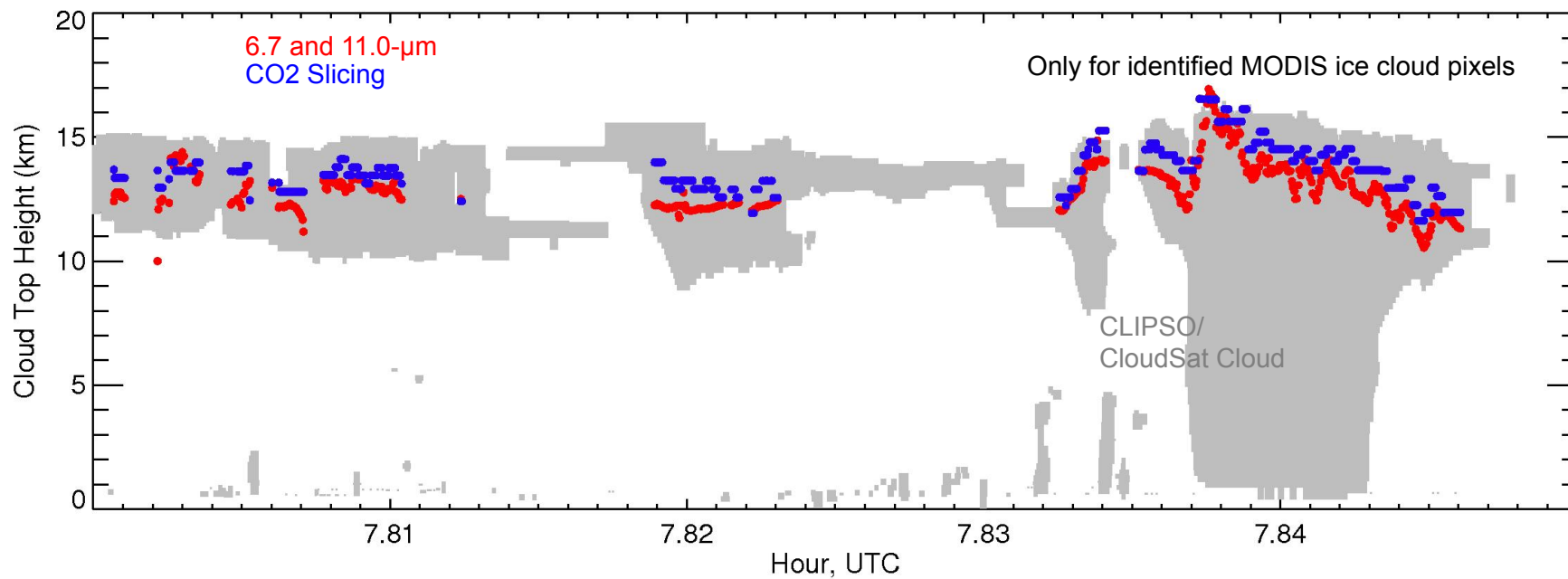
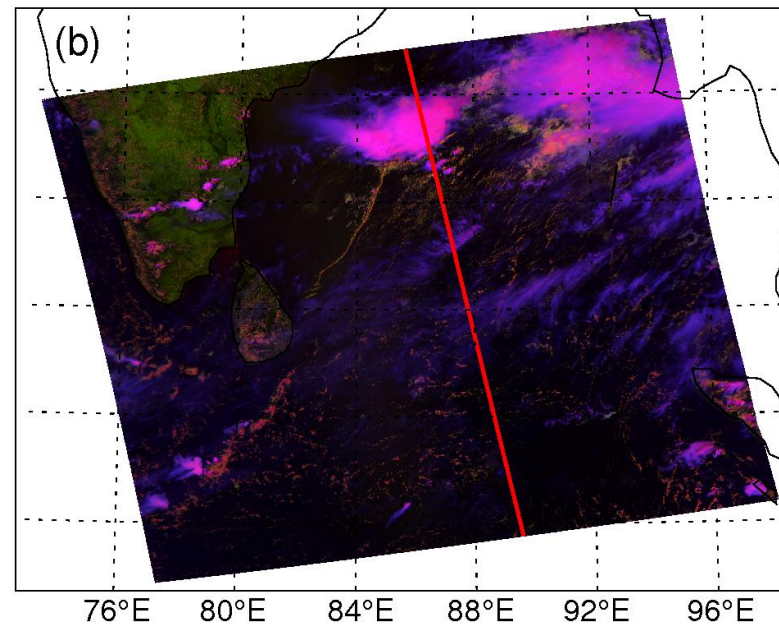
Could be less affected by variation of clear sky radiance with respect to pixel-based method.

# Ice Cloud Top Temperatures/Heights

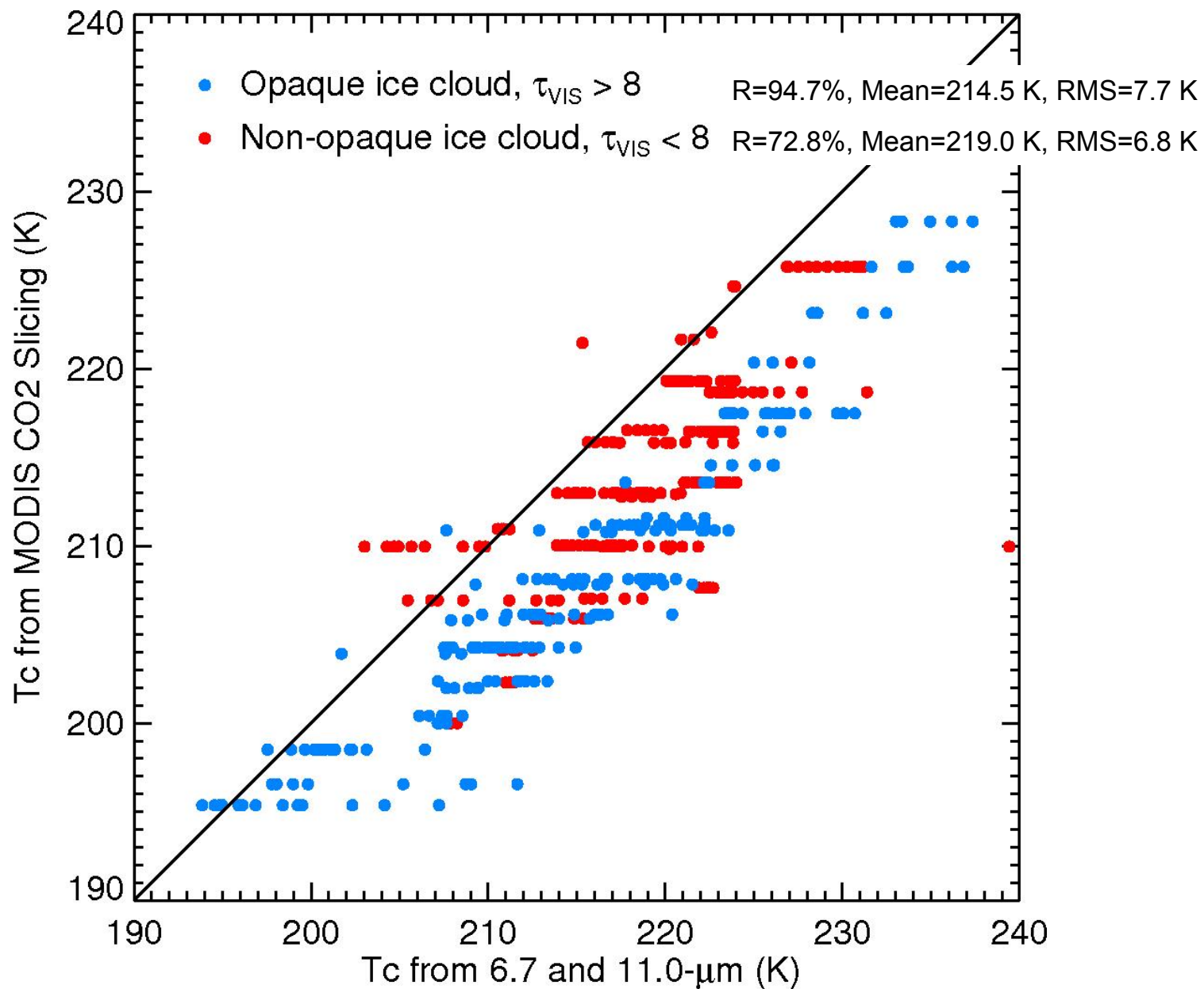
True RGB, Aqua MODIS 2007.05.22 07:45



False RGB



## Ice Cloud Top Temperatures/Heights



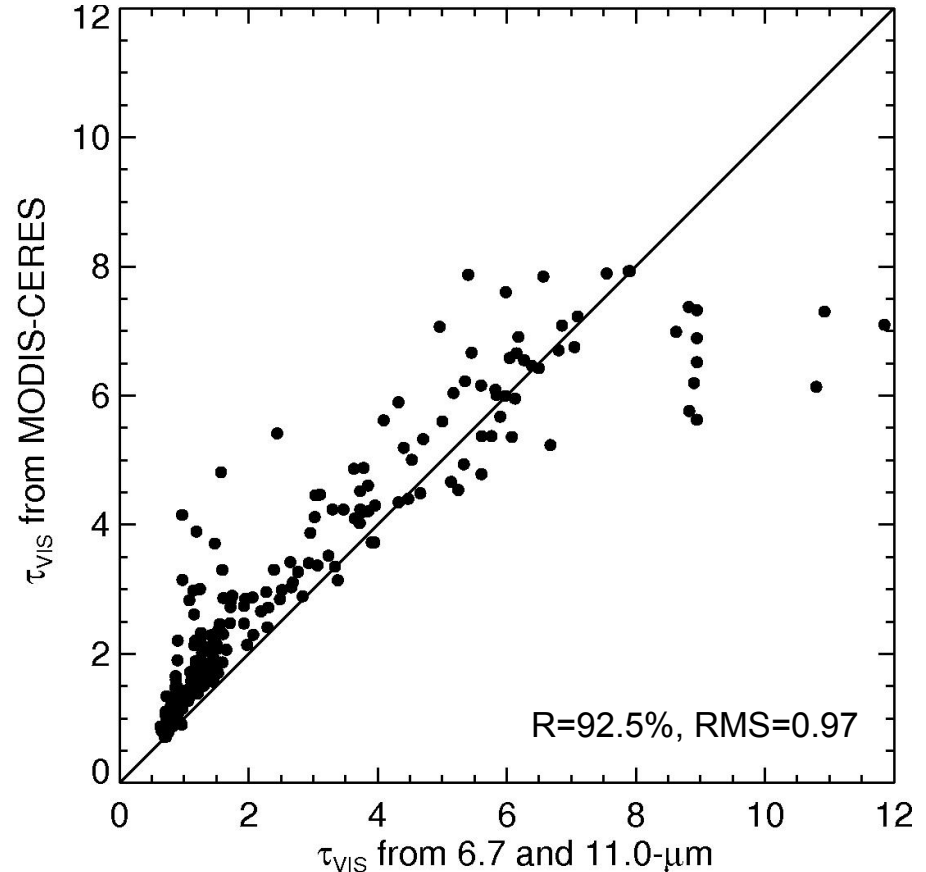
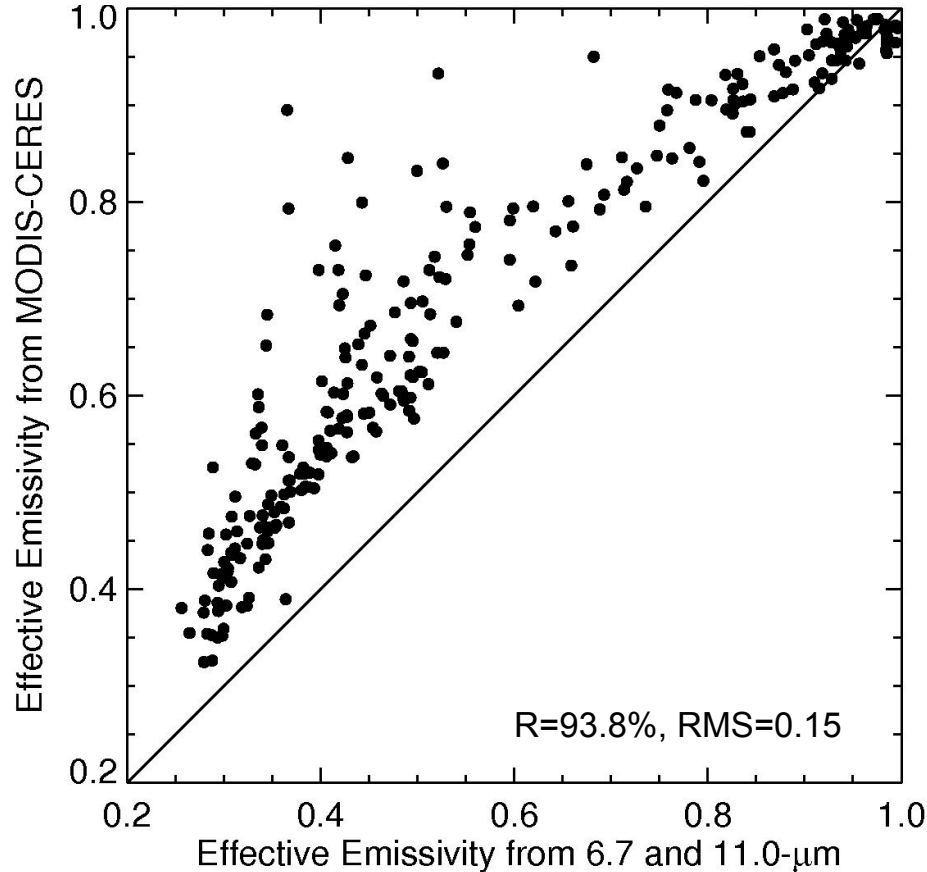
# Non-opaque Ice Cloud Effective Emissivity and Optical Thickness (1)

Once obtaining Tc for non-opaque ice clouds,  $N\epsilon$  is derived from

$$I_{11} = (1 - N\epsilon_{11}) \cdot I_{11}^{clear} + N\epsilon_{11} \cdot I_{11}^{cloud}$$

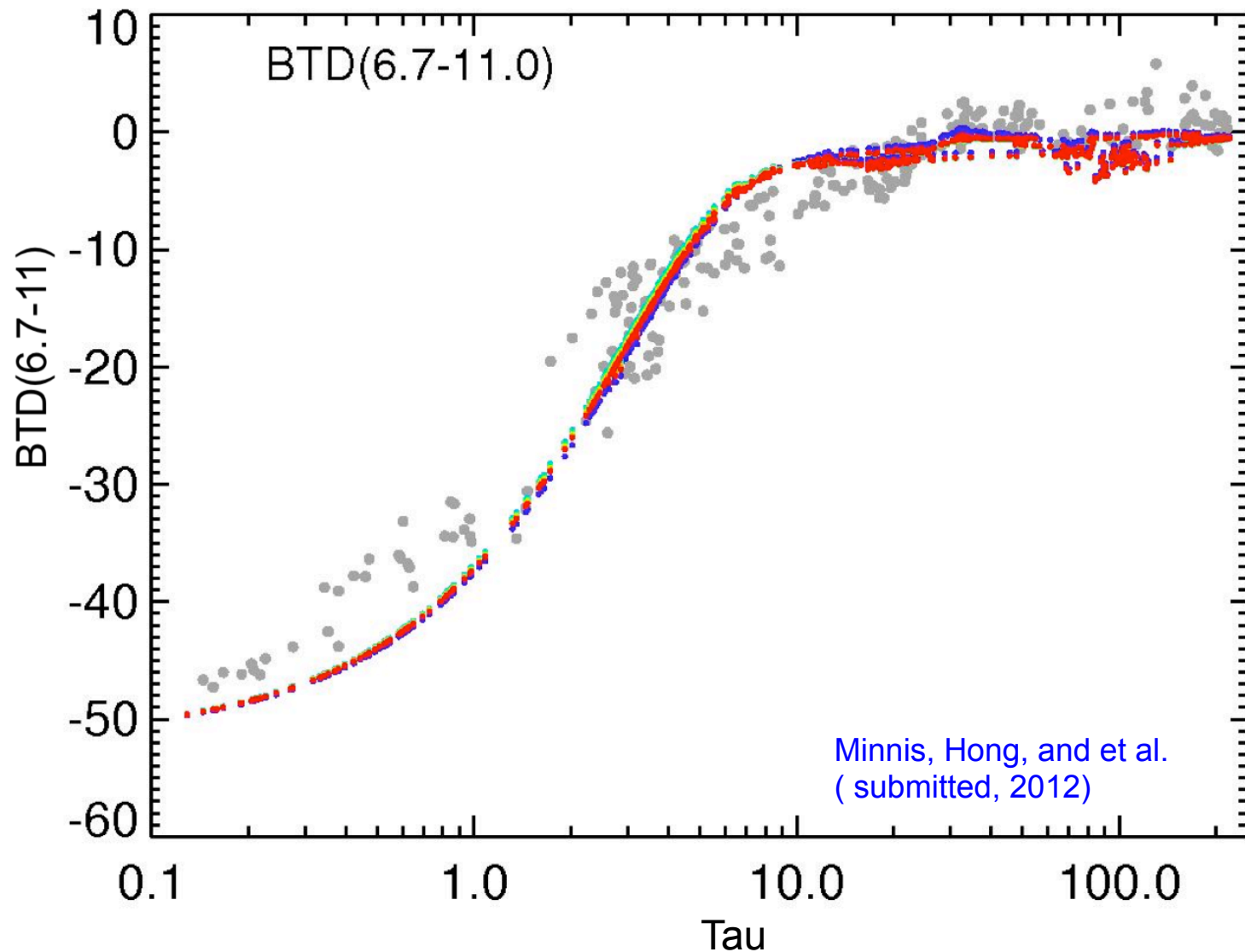
$N\epsilon$  is then converted to visible optical thickness using parameterizations from Minnis et al. (1993, 1998),

$$\tau_{VIS} = -2.13 \cdot \ln(1 - N\epsilon)$$



## Non-opaque Ice Cloud Optical Thickness from 6.7 and 11.0 $\mu\text{m}$ (2)

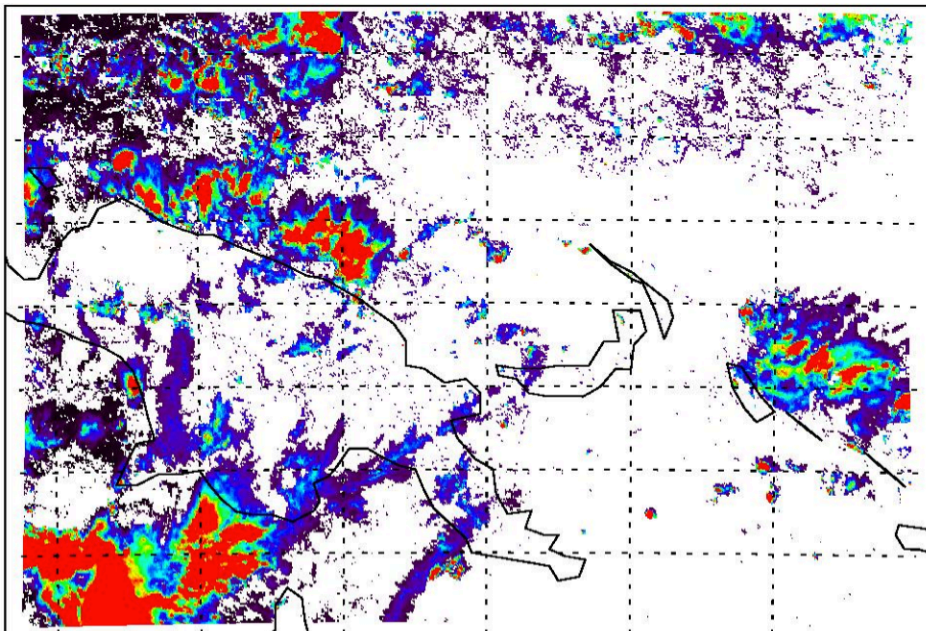
Using  $\text{BTD}(6.7-11.0)$  for  $\tau$ ,  
weakly sensitive to  $\text{Re}$



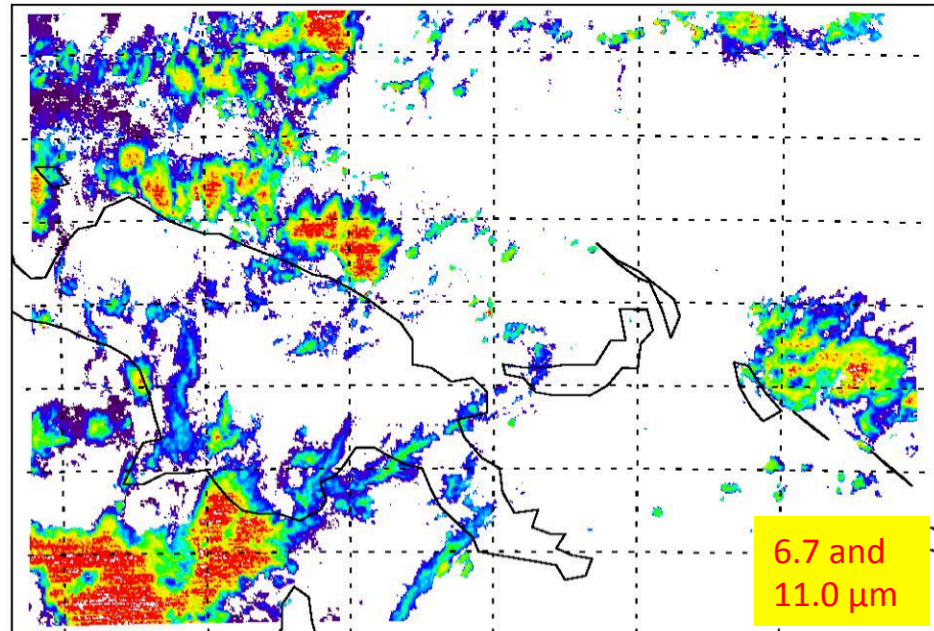
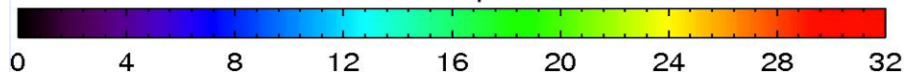


# Optical thickness from 6.7 and 11.0 $\mu\text{m}$ vs. VISST Retrievals

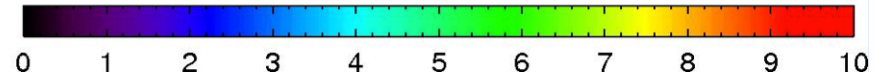
VISST - the Visible Infrared Solar-infrared Split Window Technique (Minnis et al., 1998, 2011)



VISST Ice Cloud Optical Thickness



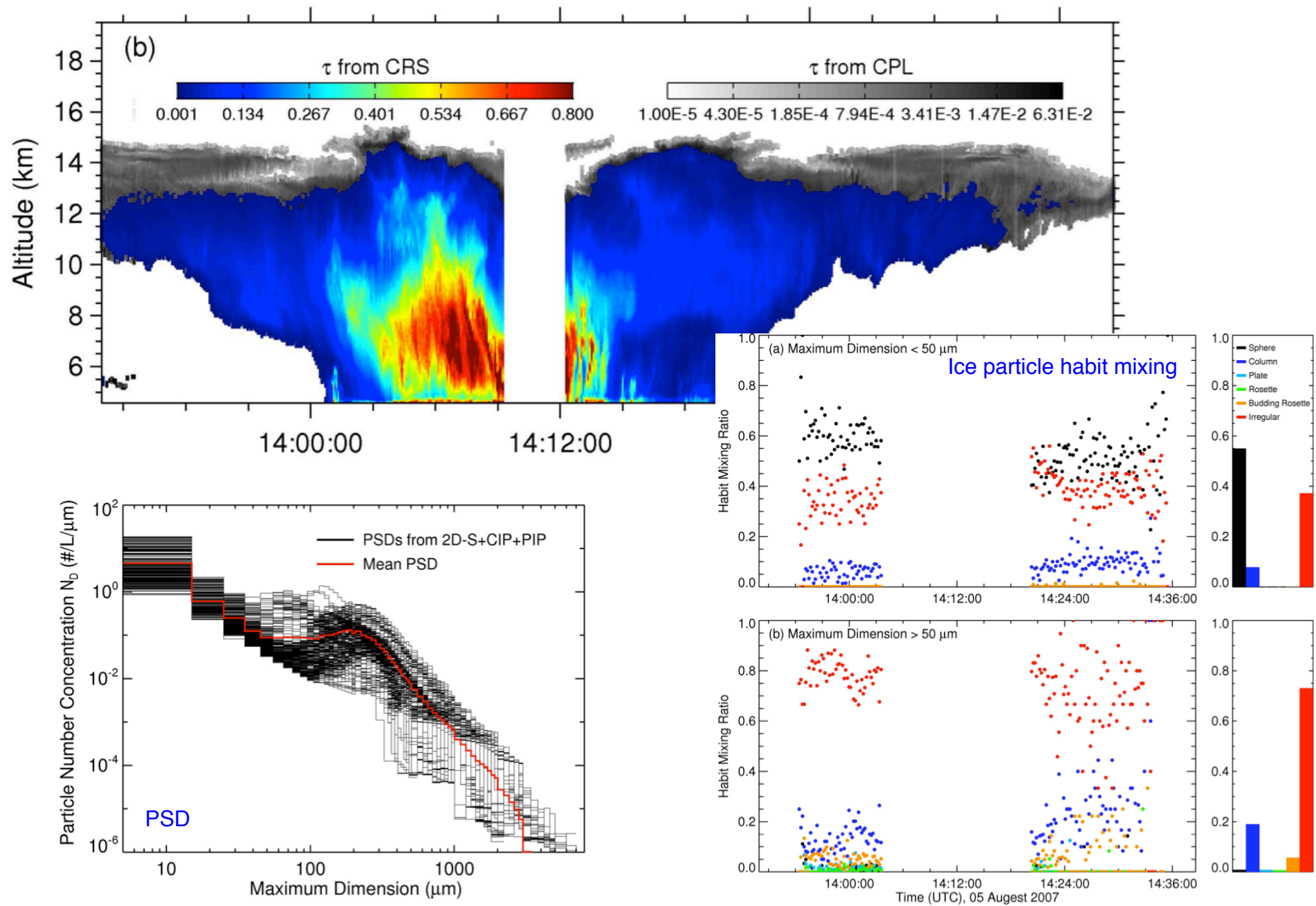
WIT Ice Cloud Optical Thickness



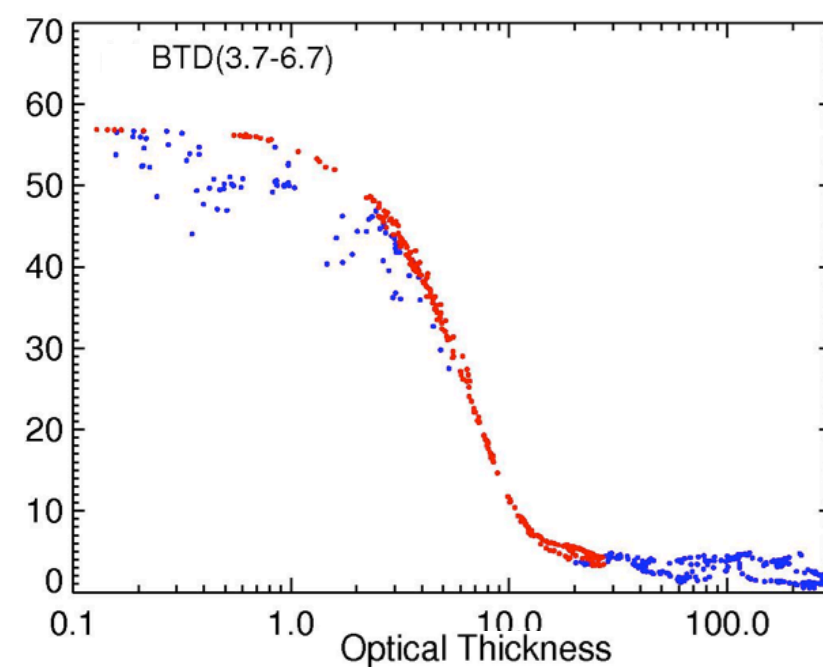
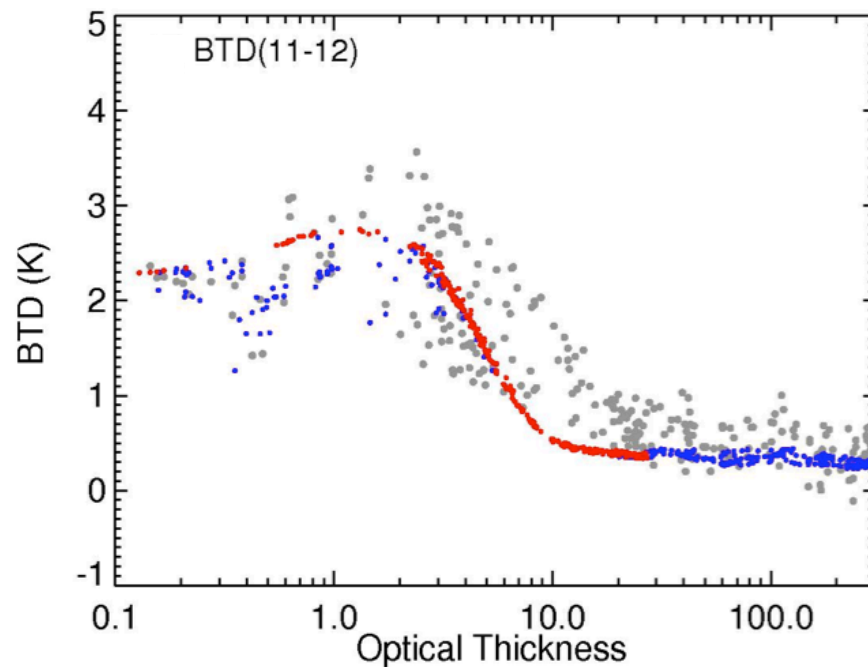
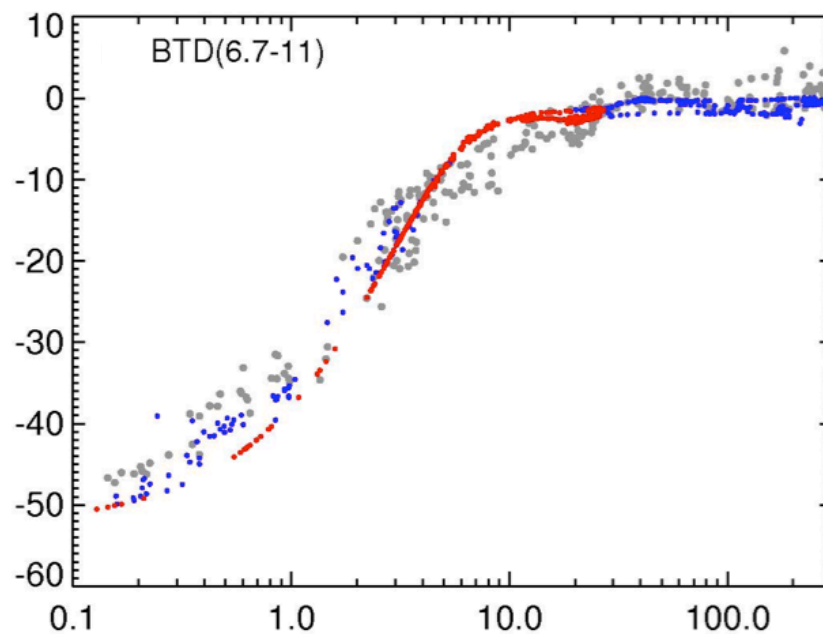
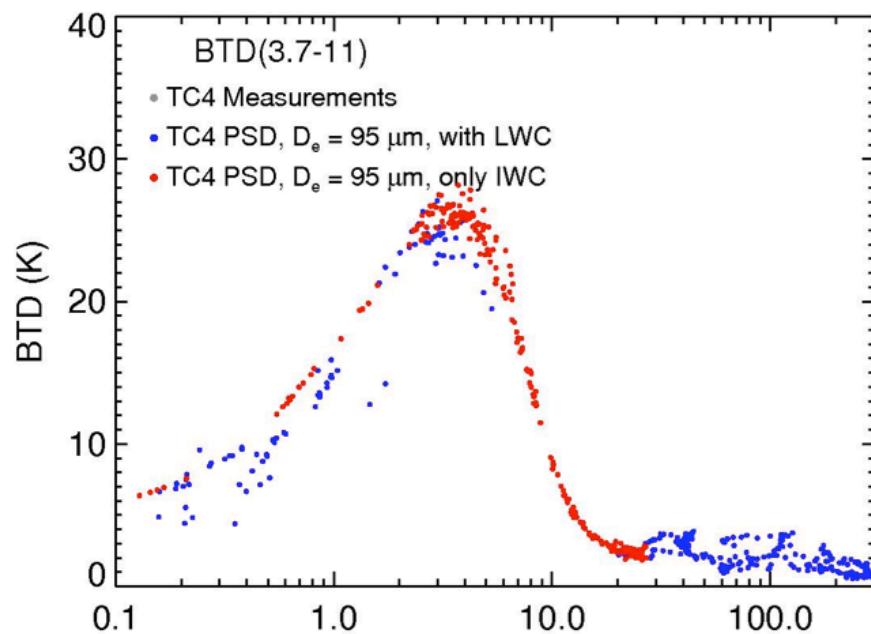
Opaque Ice Cloud Optical Thickness from IR Retrieval - Blackbody Limitation



# IR Radiances Over DCC: Observations and Simulations (Minnis, Hong, and et al., 2012)



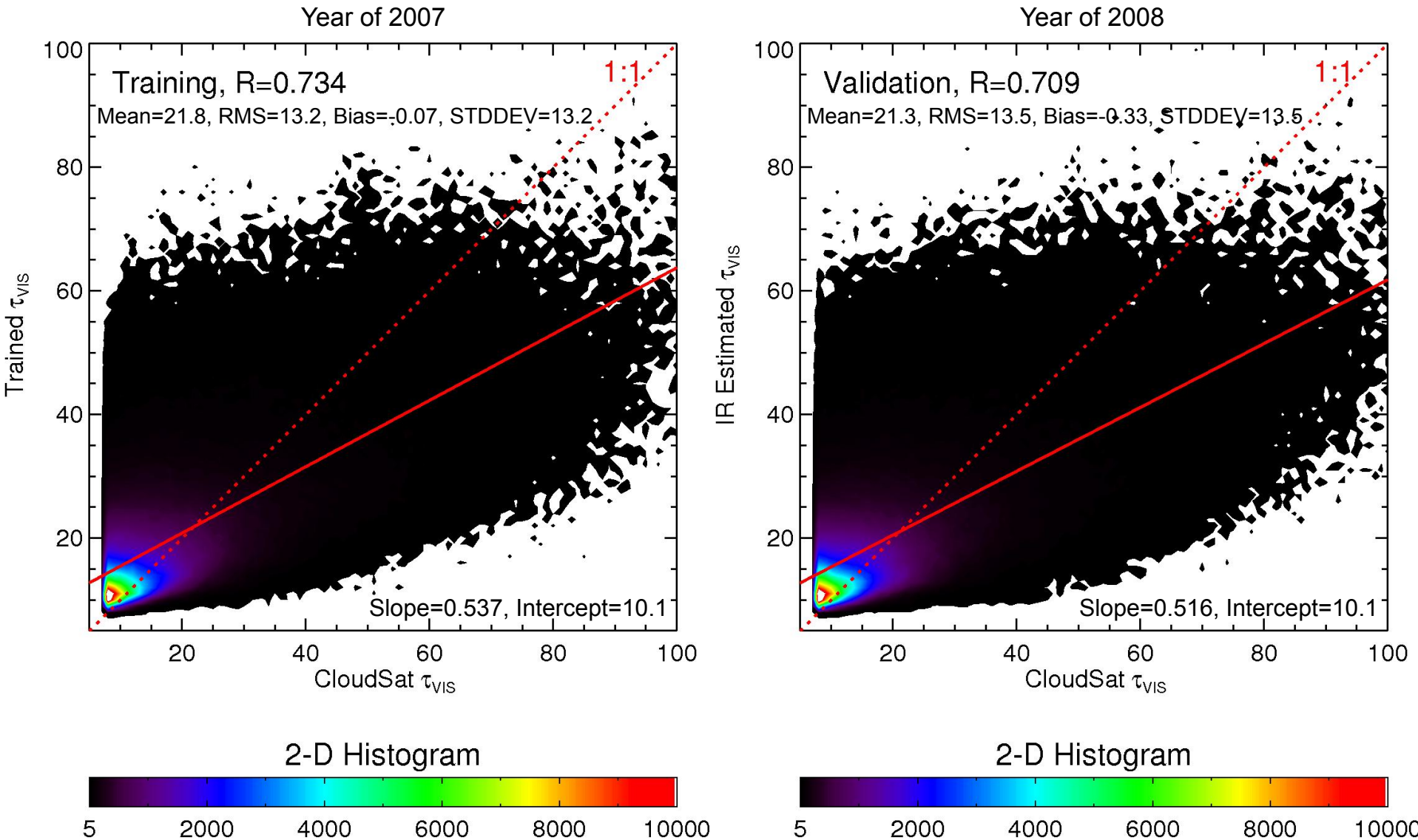
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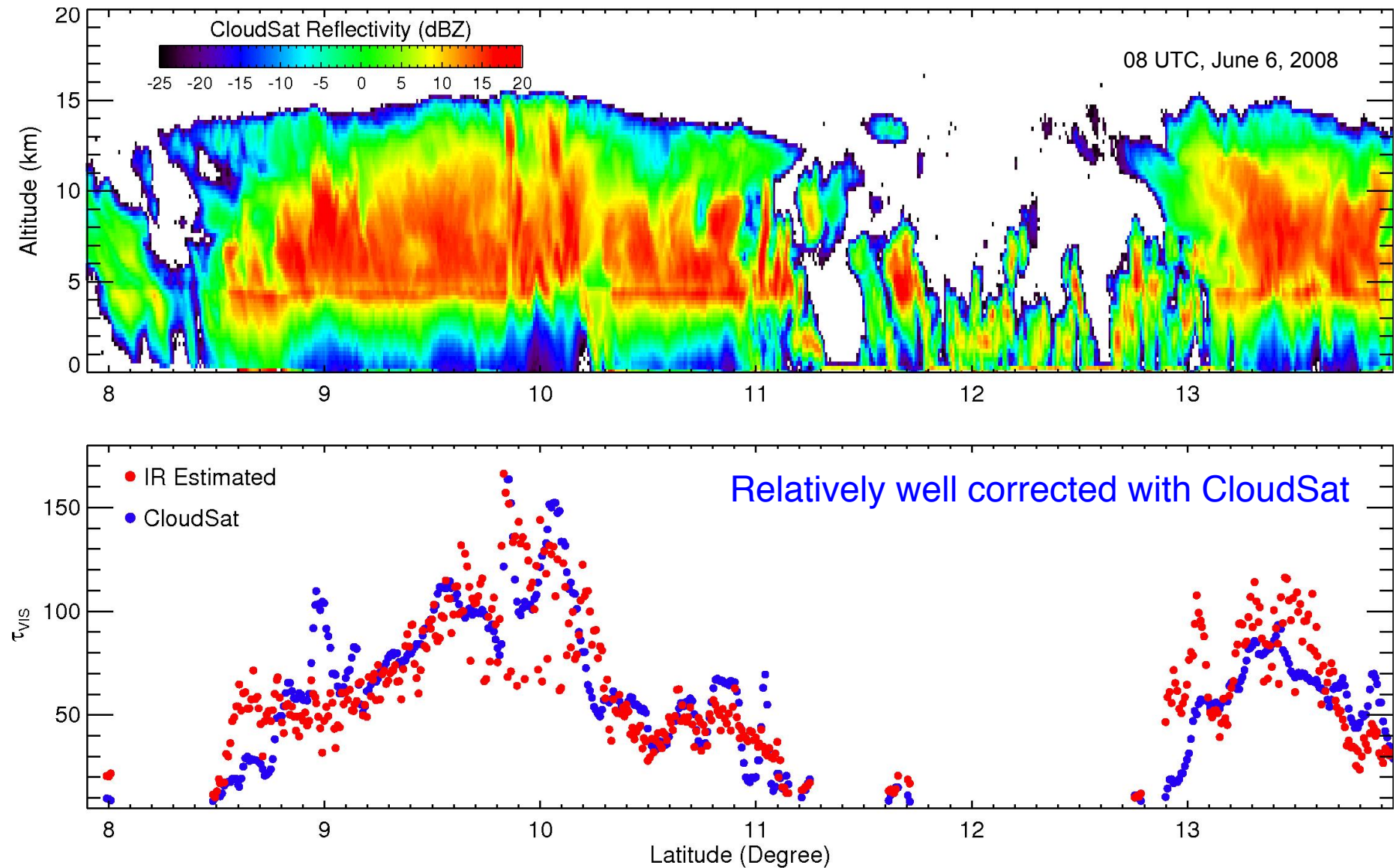
# Opaque Ice Cloud Optical Thickness from IR Measurements: Neural Network Training and Validation

**Training data:** 2007 LaRC C3M Data at night (globally) (*Kato et al., 2011*)

**Validation data:** 2008 LaRC C3M Data at night (globally)

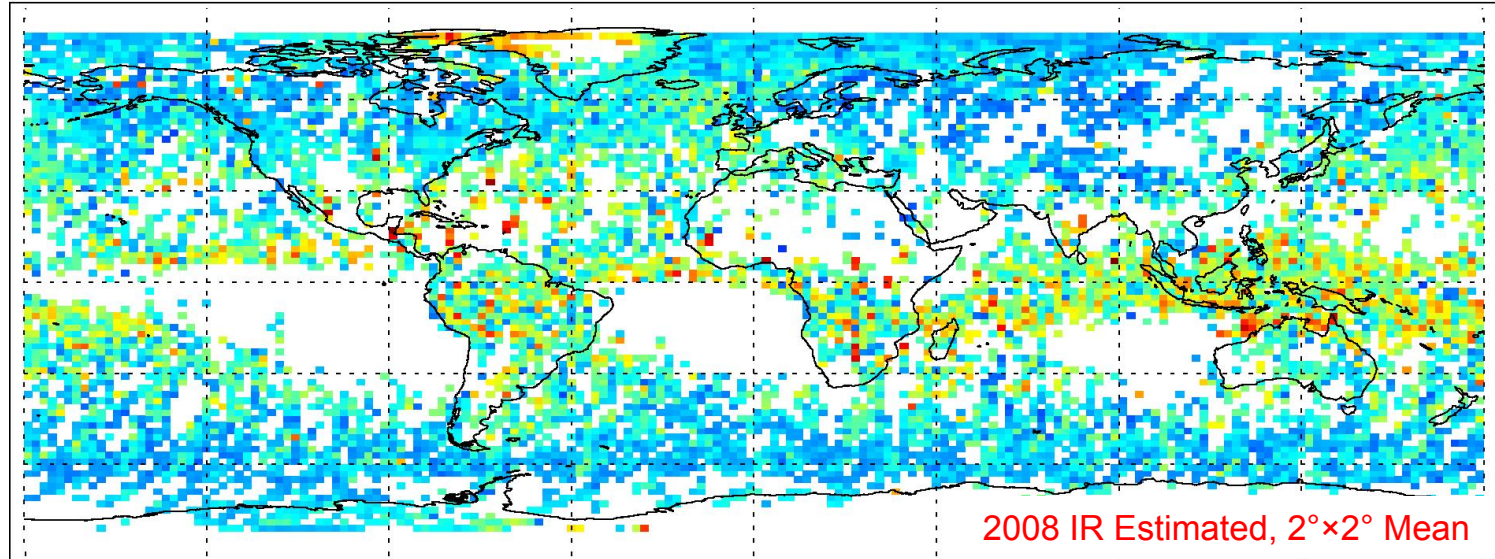
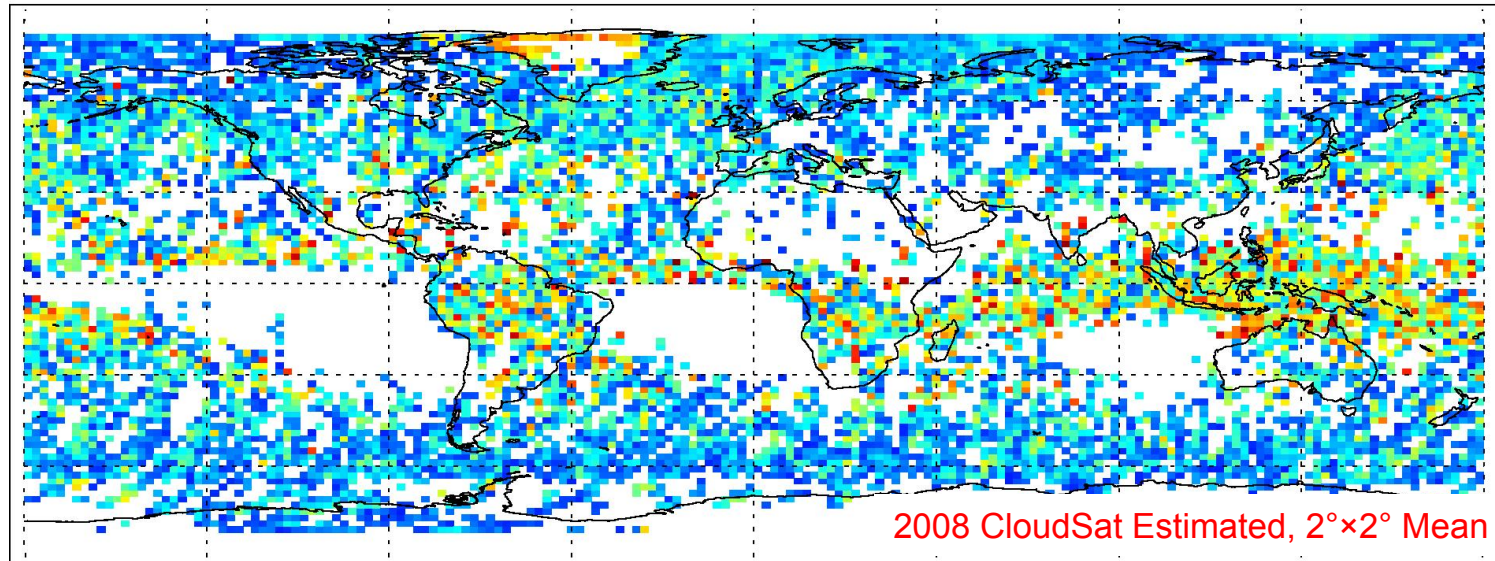


# Opaque Ice Cloud Optical Thickness from IR Measurements: An Example of Application





# Geographical Distribution of Opaque Ice Cloud Optical Thickness at Nighttime



Visible Optical Thickness





# Summary

- Water vapor band at 6.7  $\mu\text{m}$  and IR window band at 11.0  $\mu\text{m}$  are used to estimate non-opaque ice cloud-top temperature. Agree well with CO<sub>2</sub> slicing method, slightly higher.
- After obtaining T<sub>c</sub>, N<sub>ε</sub> and Tau of non-opaque ices are derived from parameterization method and physical retrieval method. Tau is very consistent with the results from visible measurements.
- Opaque ice clouds become IR blackbody emission  
Rigorous radiative transfer modeling using *in situ* measurements indicates that 3.7, 6.7, 11.0, and 12.0- $\mu\text{m}$  still show some sensitivities to opaque ice cloud, at least to Tau  $\sim$  20.
- Go beyond IR blackbody limitation for estimating opaque ice cloud Tau from a neural network method built on the basis of collocated MODIS IR measurements and CloudSat-derived Tau in 2007.  
2008 data used to validate the technique find that IR-estimated Tau agrees with CloudSat-derived Tau for opaque ice clouds with uncertainty of  $\sim$ 63%.